Adsorption of strontium and caesium onto an Na-illite and Na-illite/Na-smectite mixtures: Implementation and application of a multi-site ion-exchange model

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ABSTRACT

The application of a multi-site ion-exchange (MSIE) model to the prediction of contaminant behaviour onto soils or sediments requires a coherent adsorption database. Following the same approach exposed in preceding works on smectite (Siroux et al., Appl. Geochem. 87, 167; Wissocq et al., Appl. Geochem., 93, 167), a database is here implemented to allow modelling the adsorption of radionuclides $^{90}$Sr and $^{137}$Cs onto Na-illite, one of the major clay mineral encountered in the terrestrial environment. This database allows describing the adsorption properties of Sr$^{2+}$ and Cs$^+$ competing with major cations H$^+$ — considered as the reference cation — and Na$^+$ onto an Na-illite. To obtain the adsorption site capacities and selectivity coefficients of illite sites for Na$^+$ towards H$^+$, a saturation curve of Na$^+$ on Na-conditioned illite was acquired. Adsorption isotherms issued from literature have been reinterpreted using the MSIE model to obtain the adsorption properties of Sr$^{2+}$ and Cs$^+$ onto purified and Na-conditioned illite. Afterwards, the predictability and robustness of the database were verified simulating experimental adsorption of Sr$^{2+}$ and Cs$^+$ onto Na-illite/Na-smectite mixtures.

Keywords: cesium; caesium; strontium; sodium; illite; adsorption; modelling; multi-site ion-exchange model; additivity; clay mineral; soils
1 INTRODUCTION

The development of nuclear energy requires knowledge of potential consequences on the environment. $^{90}$Sr and $^{137}$Cs are two of the most important daughters of uranium and plutonium fission reaction (England and Rider, 1993). Both radio-isotopes were scattered in the environment after nuclear accidents such as Chernobyl and Fukushima-Daiichi, but also from the fall-out of nuclear weapons detonations.

In soils, $^{90}$Sr and $^{137}$Cs are strongly adsorbed and their behaviours are mainly dominated by their interactions with clays. As clay minerals bear structural negative surface charges, the adsorption of cation, as $\text{Sr}^{2+}$ and $\text{Cs}^+$, is facilitated. The adsorption of these radionuclides onto clay minerals, such as bentonite (Galamboš et al., 2013; Montavon et al., 2006), illite (Fuller et al., 2015; Mahoney and Langmuir, 1991), kaolinite (Chen et al., 2015; Erten et al., 1988; Reinoso-Maset and Ly, 2014) or vermiculite (Dzene et al., 2015; Long et al., 2014), has been extensively studied in the literature. Illite is especially referenced as the major adsorbent for $\text{Cs}^+$ onto strong affinity low concentration sites, referred as frayed edge sites (Brouwer et al., 1983; Cremers et al., 1988; Sawhney, 1972).

Many adsorption models were developed in order to represent the adsorption behaviours of $\text{Sr}^{2+}$ and $\text{Cs}^+$ onto clay minerals, sediment, or soils. In most of these models, adsorption is described as an ion-exchange (Robin et al., 2015) or a surface complexation (Nie et al., 2017) reaction. On the other hand, Bradbury and Baeyens (1999) have considered both adsorption and complexation phenomena, and have developed a model based on a mechanistic description of the cation adsorption between a clay minerals and a liquid phase — namely the two-sites protolysis non-electrostatic surface complexation and cation-exchange model or 2SPNE SC/CE.

These models, considering ion-exchange or surface complexation, are currently applied on more complex phase mixtures as soils. Many authors used a generalized composite approach, which consists in fully describing the complex material as a unique adsorbent, e.g. Motellier et al. (2003) or Jacquier et al. (2004). This approach allows a good representation of the adsorption properties of the studied material, but it often impedes the extrapolation of the obtained data onto different materials. Another approach considers the adsorption properties of a complex adsorbent as resulting from the additivity of simpler components. This approach requires a large amount of inputs data — description of adsorption reaction onto the main reactive mineral contained in the studied material —, but it allows a better extrapolation of the modelling once the database is completed. Missana et al. (2014b) have applied a model derived from the 2SPNE SC/CE and the additivity approach in order to simulate the adsorption of caesium onto two Spanish clays.

The multi-site ion-exchange (MSIE) formalism (Stammose et al., 1992), which will be used in this study, has been successfully applied in several studies. For instance, Tertre et al.
have modelled the Zn\(^{2+}\) behaviour onto montmorillonite and simulated successfully its adsorption onto different soils and sediments. Reinoso-Maset and Ly (2014) have studied the adsorption of major cations such as Na\(^{+}\), Ca\(^{2+}\), K\(^{+}\), and Mg\(^{2+}\) onto pure kaolinite, and the behaviour of Cs\(^{+}\) at trace level. Finally, the studies carrying out the adsorption of Sr\(^{2+}\) and Cs\(^{+}\) onto an Na-smectite (Siroux et al., 2017), and onto Ca-illite and -smectite (Wissocq et al., 2018), served as a basis for a general adsorption database for the MSIE formalism.

Applying the additivity approach in the framework of the MSIE model enables simulating adsorption onto complex phases. The Sr\(^{2+}\) and Cs\(^{+}\) adsorption onto sedimentary rock (Wissocq et al., 2018) and their mobilities in a clayey rock (Savoye et al., 2012; Savoye et al., 2015) were also successfully simulated using this approach.

The aims of this paper are: i) implementing the database for the MSIE model with the adsorption properties of Na\(^{+}\), H\(^{+}\), Cs\(^{+}\) and Sr\(^{2+}\) onto Na-illite; and ii) verifying the additivity approach on Na-illite/Na-smectite mixtures using the MSIE model. First, a purified Na-conditioned illite du Puy will be described in terms of site capacity (SC) and corrected — of adsorbed species activity coefficients — selectivity coefficients for the Na\(^{+}\)/H\(^{+}\) exchange (K\(^{i}_{\text{Na}^{+}/\text{H}^{+}}\)) for each site. Second, adsorption isotherms of Sr\(^{2+}\) and Cs\(^{+}\) onto purified and Na-conditioned illite samples will be compiled from literature and modelled in order to determine the selectivity coefficients for the M\(^{n+}\)/nH\(^{+}\) exchanges (K\(^{ii}_{\text{M}^{n+}/n\text{H}^{+}}\)). Finally, the implemented database, including the parameters from Siroux et al. (2017) for Na-smectite, will be used to predictively simulate experimental adsorption isotherms of Sr\(^{2+}\) and Cs\(^{+}\) onto illite/smectite mixtures.

2 MATERIALS AND METHODS

2.1 Materials

A SWy-2 smectite purchased from the Clay Mineral Society, and an illite du Puy were used in this study. These clay minerals were purified in order to isolate the montmorillonite (SWy-2) and the illite fractions from these clay materials. The applied purification protocol consists in eliminating organic matter as much as possible with H\(_2\)O\(_2\) at 80°C, and dissolving calcite with a 0.1 mol.kg\(^{-1}\) HCl solution. The resulting smectite and illite solids were dispersed in pure water (solid/solution ratio at 1:10) and sieved at 63 and 25 µm, respectively. This protocol is further detailed in Gaucher (1998), Peynet (2003), and Reinoso-Maset and Ly (2014). Both clay minerals were then Na-conditioned by five successive washings with a 1 mol per kg of water (mol.kg\(^{-1}\)) NaCl, and a final washing with a 0.1 mol.kg\(^{-1}\) NaCl solution. Finally, the Na excess was removed by rinsing with Milli-Q water (Millipore, 18.2 MΩ.cm\(^{-2}\)), then with ethanol. Purification and conditioning treatments for the clay minerals were confirmed by XRD analysis. The purified SWy-2 still presented traces of well-crystalized quartz and no significant trace of illite, while illite presented traces of quartz and kaolinite.
All chemicals were at least of analytical grade. The stable Sr solutions was prepared with a SrCl₂·6H₂O salt (99.995% trace metals basis, Sigma Aldrich), while CsCl (99.999% trace metals basis, Sigma Aldrich) was used for the stable Cs solutions. Radioisotopes ⁸⁵Sr, ¹³⁷Cs, and ²²Na were provided by Cerca LEA. Ionic strength was fixed using NaCl (Analar Normapur, VWR).

2.2 EXPERIMENTAL PROTOCOL

2.2.1 Analytical methods

The aqueous concentrations measurements of Sr, Na, Ca, K, and Mg were determined by inductively coupled plasma atomic emission spectroscopy (ICP-AES, Activa, Horiba Jobin Yvon), diluting aliquots with 2% sub-boiling HNO₃ (DuoPur, Milestone). Measurements were made using calibrations with five points at 407.771 nm for strontium, 589.592 nm for sodium, 317.933 nm for calcium, 766.490 nm for potassium, 273.553 nm for magnesium. The aqueous concentrations of Cs were measured by inductively coupled plasma mass spectrometry (ICP-MS, Element 2, Thermo Scientific). The pH values were measured with a combined microelectrode (pH InLab, Mettler Toledo) by redetermination of the potential/pH plot of the electrode and by calculating the value of the solutions from their potentials (buffer solutions 4.01, 7.00, 9.18, Merck). Uncertainties from the linear regression are typically in the range of 0.05 pH-units.

2.2.2 Characterization of clay minerals

The cation-exchange capacity (CEC) and the exchangeable cations measurement of the purified smectite and illite were measured as in Siroux et al. (2017). The CEC measurement consists in extracting all exchangeable cations from the clay mineral with a 0.1 mol.kg⁻¹ KCl solution with addition of KOH to reach a pH ca. 10. Then, adsorbed K⁺ was extracted with a 0.1 mol.kg⁻¹ CsCl solution then measured by ICP-AES (Table 1). The sum of the exchangeable cations is close to the measured CEC. As expected, sodium is the major exchangeable cation in purified clay minerals but there are still traces of K⁺ and Mg²⁺.
Table 1. Exchangeable cations concentrations and CEC measurement of purified Na-clay minerals.

<table>
<thead>
<tr>
<th>Sample</th>
<th>Na (meq.100g⁻¹)</th>
<th>Ca (meq.100g⁻¹)</th>
<th>Mg (meq.100g⁻¹)</th>
<th>K (meq.100g⁻¹)</th>
<th>Sum of exchangeable cations (meq.100g⁻¹)</th>
<th>CEC (meq.100g⁻¹)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Swy-2 smectite</td>
<td>79</td>
<td>&lt;l.o.d.</td>
<td>0.4</td>
<td>2.5</td>
<td>81.9</td>
<td>83.3</td>
</tr>
<tr>
<td>Illite du Puy</td>
<td>19.6</td>
<td>&lt;l.o.d.</td>
<td>&lt;l.o.d.</td>
<td>1.1</td>
<td>20.7</td>
<td>20.9</td>
</tr>
</tbody>
</table>

l.o.d: limit of detection

2.2.3 Na-saturation curve and adsorption isotherm onto pure clay mineral mixtures

Experiments were done by batch method at room temperature (22 ± 5 °C) in 10 mL polycarbonate Nalgene centrifuge tubes (3138-0010, Thermo Scientific). All solutions and suspensions were made by weighing. All equilibration and agitation of the suspensions were done using an orbital shaker.

Saturation curve of Na⁺ onto pure Na-illite was made with a 19 g.L⁻¹ S/L ratio. A 0.1 g of illite aliquot was dispersed in 5 mL of 0.02 mol.kg⁻¹ NaCl and equilibrated during 24 h. Solutions of 1, 0.1, or 0.01 mol.kg⁻¹ of either HCl or NaOH were used in order to adjust the pH of the suspensions to values ranged between 3 and 11. Suspensions were then equilibrated during ca. 24 h, and subsequently spiked with ²²Na — 300 µL of NaCl in HCl 0.1 N added containing 13 kBq/mL ²²Na — and equilibrated during 7 days. Suspensions were ultra-centrifuged at 33,000 g during 30 min (Beckman Optima LE-80K, Beckman Ultracentrifuge); 1 mL aliquots of the supernatants were collected, and activity was counted by gamma counting (1480 Wizard 3, Perkin Elmer). Total Na content (stable and radioisotope) was analysed by ionic chromatography (standard Chem-Lab).

The partition coefficients \( R_D \) (L.kg⁻¹) were calculated using,

\[
R_D = \left( \frac{C_0}{C_{eq}} - 1 \right) \frac{V}{m}
\]

(1)

where \( C_0 \) and \( C_{eq} \) are respectively the initial and equilibrium activities of Na, Sr, or Cs, \( V \) is the total volume of solution (L), and \( m \) is the total dry mass of solid (kg).

The initial activities were determined using a blank procedure, i.e. using the same protocol without clay minerals in batches. The adsorbed concentrations were then calculated from the initial concentrations and the partition coefficient.

Pure clay mixtures experiments were prepared by weighing, and dry mechanic stirring the dried pure minerals. In these experiments, a 3.8 g.L⁻¹ solid/solution ratio was used.
Illite/smectite mixtures were first equilibrated with a 0.1 mol.kg\(^{-1}\) NaCl electrolyte for 24 hours. The pH was adjusted ca. 7 with 1, 0.1, or 0.01 mol.kg\(^{-1}\) solutions of either HCl or NaOH, and suspensions were left under agitation for ca. 24 h. The concentration of stable caesium — from \(10^{-9}\) to \(10^{-2}\) mol.kg\(^{-1}\) in CsCl — or strontium — from \(10^{-6}\) to \(10^{-2}\) mol.kg\(^{-1}\) in SrCl\(_2\) — was added in the centrifuge tubes. Suspensions were also spiked with \(^{85}\)Sr or \(^{137}\)Cs (ca. 3,000 Bq), then left under agitation during two and five days for Sr and Cs, respectively. Finally, the suspensions were ultra-centrifuged at 90,000 \(g\) during 30 min. (Beckman Optima LE-80R, Beckman Ultracentrifuge); 1 mL aliquots of the supernatants were collected, and activity was counted by gamma counting. The initial concentrations of Sr were determined in ICP-AES, and the Cs concentrations in ICP-MS.

2.2.4 Verification of the additivity approach

Purified and Na-conditioned SWy-2 smectite and illite du Puy were used for these experiments. Two mixtures were prepared by mechanical mixing powder of purified clay minerals. The exact percentages of each clay mineral mixture are detailed in Table 2. For M1 and M2 mixtures, adsorption isotherms of Cs\(^+\) and Sr\(^{2+}\) vs. concentration were acquired at pH 7 and \(I = 0.1\) mol.kg\(^{-1}\) NaCl.

<table>
<thead>
<tr>
<th>Sample</th>
<th>% illite</th>
<th>% smectite</th>
</tr>
</thead>
<tbody>
<tr>
<td>M1</td>
<td>30.2</td>
<td>69.8</td>
</tr>
<tr>
<td>M2</td>
<td>70.1</td>
<td>29.9</td>
</tr>
</tbody>
</table>

2.3 Modelling

Isotherms were modelled using the MSIE model, which is fully detailed elsewhere (Jacquier et al., 2004; Motellier et al., 2003; Reinoso-Maset and Ly, 2014; Savoye et al., 2012; Savoye et al., 2015; Siroux et al., 2017; Stammose et al., 1992; Tertre et al., 2009; Wissocq et al., 2018). Briefly, in this model clay minerals are defined as exchangers with different charged adsorption sites (\(X_i^-\)), which are expressed in terms of site capacities (SC\(_i\)) and corrected — of adsorbed species activity coefficients — selectivity coefficients for a Na\(^+\)/H\(^+\) exchange (\(K_{Na^+/H^+}^{X_i^-}\)). Sites capacities are obtained from the Na-saturation isotherm describing the cation saturation of the homo-ionic exchanger vs. pH. The equilibrium between Na\(^+\) and H\(^+\) on negatively charge type of site \(X_i^-\) is expressed by the following equilibrium reaction.

\[
X_i^- - Na^+ + H^+ ⇋ X_i^- - H^+ + Na^+
\]

The equilibrium constant of this equation can be written as,
\[ K_{Na^+/H^+}^X_i = \frac{[X_i^- - H^+].[Na^+] \times \gamma_{Na^+}}{[X_i^- - Na^+].[H^+] \times \gamma_{H^+}} \times \left( \frac{f_{X_i^- - H^+}}{f_{X_i^- - Na^+}} \right) \] (2),

where \([\ ]\) is the solution (mol.kg\(^{-1}\)) concentration, \(\gamma\) is the adsorbed concentration (mol.kg\(^{-1}\) of solid) of species, \(\gamma\) are the activity coefficients of aqueous species — calculated here using Davies (1962) equation —, and \(f\) are activity coefficients of adsorbed species.

Since the activity coefficients of adsorbed species \(f\) are not known, it is assumed that the ratio between these coefficients is constant. Thus, the apparent constants — called corrected selectivity coefficients \(K_{Na^+/H^+}^X_i\) — are expressed as follows.

\[ K_{Na^+/H^+}^X_i = \frac{[X_i^- - H^+].[Na^+]}{[X_i^- - Na^+].[H^+]} \times \frac{\gamma_{Na^+}}{\gamma_{H^+}} \] (3)

It is worthy to notice that the exchanger is supposed not to undergo any phase modification when a physicochemical parameter of the solution is changed, e.g. pH, ionic strength, temperature...

Sites \(i\) are also described by their site capacities \(SC_i\) in mol.kg\(^{-1}\) of solid, which equals to the sum of the adsorbed species concentrations.

\[ SC_i = [X_i^- - H^+] + [X_i^- - Na^+] \] (4)

In the present study, sites capacities are obtained from the Na-saturation curve describing the cation saturation of the homo-ionic exchanger, express in term of adsorbed species concentration (mol/kg of solid) vs. pH. Each site exchange capacity is expressed by a plateau corresponding to an Na\(^+\)/H\(^+\) exchange reaction.

The corrected selectivity coefficients associated to the trace element (Sr\(^{2+}\) and Cs\(^+\)) exchanged with Na\(^+\) are determined fitting the adsorption isotherms of these trace elements onto homo-ionic Na-conditioned exchanger vs. the solution pH or the equilibrium concentration of the trace element in solution. The corrected selectivity coefficient for the M\(^{m+}\)/Na\(^+\) exchange is written as follows.

\[ K_{mNa^+/M^{m+}}^X_i = \frac{[X_i^- - M^{m+}]_m.[Na^+]_m.\gamma_{Na^+}^m}{[X_i^- - Na^+]_m.[M^{m+}].\gamma_{M^{m+}}} \] (5)

Calculation of \([X_i^- - M^{m+}]_m\) permits calculating experimental partition coefficient \(R_D\) using the following expression:

\[ R_D = \frac{[M]}{[M]} = \frac{\sum_i([X_i^- - M^{m+}])}{[M]} \] (6).
In case of illite/smectite mixtures, the total adsorption of $M^{m+}$ arises from each mass fraction contribution as follows.

$$ R_D = \frac{\sum_i [(X_i)_m - M^{m+}]_{\text{illite}} + \sum_j [(Y_j)_m - M^{m+}]_{\text{smectite}}}{[M]} $$

(7)

Uncertainties on $R_D$ are determined by propagation of errors. However, due to its low value, $R_D$ error bar expressed by the $R_D$ standard deviation calculated on three replicates was preferred — estimated to 0.15 log unit to compare the experimental data.

3 RESULTS AND DISCUSSION

3.1 PRELIMINARY REMARKS

As stated in Siroux et al. (2017) for Na-smectite, illite was supposed to be fully Na-homoionic. Despite their possible interactions in the adsorption processes, other exchangeable cations — e.g. K+ present in the exchanger — were neglected to interpret the adsorption results. Their trace level concentrations were assumed sufficiently low not to interfere with Cs+ and Sr2+. Furthermore, it is known that pH < 4 or pH > 10 the dissolution of clay minerals can be significant (Amram and Ganor, 2005; Bradbury and Baeyens, 1997), and can lead to a release of aluminium, silicon, iron, and other structural cations as K+ and Mg2+. The release of these ions — mainly $\text{Al(OH)}_{3-n}^n$ or $\text{Fe(OH)}_{3-n}^n$ — could compete with the adsorption of Sr2+ and Cs+ onto illite. The account of these phenomena is very difficult to manage, see e.g. Bradbury and Baeyens (2005). One must not forget that these phenomena are intrinsically included into the corrected selectivity coefficients. However, Al and Fe in compensating charge, as enhanced surface charges due to dissolution, may generate erratic $K_D$ values at extreme pH. In our experiments, $K_D$ value are coherent with a global trend of adsorption. Moreover, for this study, dissolutions tests performed on illite revealed a silicon lost estimated below 0.5% at pH 3 under our experimental conditions, i.e. $m/V = 19 \text{ g.L}^{-1}$, 15 days of contact time. Therefore, the dissolution effect on adsorption properties have been neglected.

3.2 DATABASE IMPLEMENTING

The first part of this study consists in implementing the database originally built from the works of Siroux et al. (2017) and Wissocq et al. (2018) with the adsorption behaviour of Sr2+ and Cs+ onto a Na-illite. This database was built from experimental data available in the literature that follows several criteria: i) the illite must be purified; ii) the illite must be Na-conditioned; and iii) adsorption isotherms must be acquired with Na+ electrolyte (NaCl or NaClO4). In case of saturation isotherm, Na+ adsorption isotherms have to attain the CEC of illite as close as possible. As it will be discussed in the following, in order to respect this latter criterion, a new saturation curve was acquired.
In some cases the original experimental data were directly extracted from Tables of the paper or manuscript, or their retrieval was made using WinDig (1996) data digitizer software. Once the data were retrieved, the MSIE modelling was applied in order to determine the selectivity coefficients by fitting the model to the data with the MS Excel Solver. Fitting uncertainties were finally obtained using the SolverAid macro from De Levie (2004).

3.2.1 Saturation of illite with Na⁺

As already detailed elsewhere (Siroux et al., 2017; Wissocq et al., 2018), the capacity of major adsorption sites is generally deduced from saturation curves expressed in terms of the concentration of adsorbed species (mol/kg of solid) over a wide range of pH values (2 to 11). These curves are acquired for each major cation of interest, onto homo-ionic clay minerals, and under chemical conditions leading to the saturation of site capacities. Each site exchange capacity is expressed by a plateau, corresponding to an exchange reaction at equilibrium between cation and proton, which determines the site capacity CEi. So, illite must be characterized in terms of SCi and KNa⁺/H⁺ studying the saturation of an Na-illite by Na⁺ vs. pH. Ly et al. (1991) and Gorgeon (1994) have previously studied the Na⁺ saturation onto illite du Puy at [Na] = 0.02 mol.kg⁻¹: the results are recalled in Fig. 1. It appears that the saturation curve of these authors increases with pH, and then levels-off to 16.7 meq.100 g⁻¹ at pH 11.5. It seems that Na⁺ did not fully saturate the illite — CEC measured at 0.248 eq.kg⁻¹ on a < 80 µm fraction by Gorgeon (1994).

As our illite du Puy sample has been purified differently from Gorgeon (1994), new experimental points obtained in this study are also shown in Fig. 1. At acidic pH, the curve follows those from Gorgeon (1994) but instead of levelling off, the experimental points continue increasing until [Na]ads = 22 meq.100 g⁻¹. Contrary to Gorgeon (1994), sodium seems to saturate the Na-illite and achieves a concentration close to the CEC.
Fig. 1. Saturation curve of Na-illite (squares) vs. pH at \( I = 0.02 \) \( \text{mol.kg}_w^{-1} \) (NaCl) and \( S/L \) ratio = 19 g.L\(^{-1} \), and representation of the modelling and the influence of the different fixation sites, compared to the results from Gorgeon (1994) (filled diamonds) and Ly et al. (1991) — empty diamonds digitized from Gorgeon (1994).

The site capacities determined by Gorgeon (1994) and Wissocq et al. (2018) — who have implemented the database for the adsorption of Cs\(^+\) and Sr\(^{2+}\) onto a Ca-illite within the same MSIE formalism — coincide perfectly with the present new Na-saturation curve. So, for the sake of consistency the SCi determined in the literature were used. This new saturation curve permits to adjust the selectivity coefficient more precisely, as the total CEC is now saturated. The SCi and new \( \log_{10} K_{Na^+}^{X_i} / H^+ \) are recalled in Table 3. The modelling satisfactorily follows the evolution of the experimental results.
3.2.2 Adsorption isotherms of trace elements

The only studies that reach the selection criteria — cf. § 3.2 — are the adsorption isotherms of Sr$^{2+}$ and Cs$^+$ onto Na-illite from Poinssot et al. (1999a, 1999b). These authors provided pH-isotherms at two different ionic strengths and a concentration isotherm for each cation. Reinterpretation of these data in the framework of the MSIE is provided.

### 3.2.2.1 Caesium adsorption

For the sake of consistency, the selectivity coefficients for the Cs$^+$/H$^+$ exchange were determined by a simultaneous adjustment of the pH- and concentration-isotherms from...
Poinssot et al. (1999a, 1999b). As already explained in Siroux et al. (2017), Gorgeon (1994) and Poinssot et al. (1999a, 1999b) also reported trace Cs⁺ adsorption onto a pure Na-saturated illite du Puy in 1 mol.kg⁻¹ NaCl and 1 M NaClO₄, respectively. These data will not be included in our discussion because the ionic strength is well outside the validity domain of the Davies (1962) equation that is used within the adsorption database — see Siroux et al. (2017) and Wissocq et al. (2018), and references therein.

The adsorption isotherms of Cs⁺ vs. pH at ionic strength 0.1 and 0.01 M NaClO₄ — i.e. 0.101 and 0.010 mol.kg⁻¹ (Novotny and Sohnel, 1988) —, their modelling, and the contribution of the different exchange sites are shown in Fig. 2. The adsorption of Cs⁺ seems to depend weakly upon pH values. From very acidic pH, R₀ values increase steadily then reach a plateau at pH ca. 4. The modelling represents well the experimental data for the pH-adsorption isotherms at I = 0.1 M (Fig. 2a), and only three out of the four sites are necessary: the X₂ site contributes weakly to the adsorption of Cs⁺ — the contribution to the total R₀ value is lower than 10⁻⁴ L.kg⁻¹. In view of the modelling, it is also possible to state that the X₀ site drives the adsorption of Cs⁺.

Fig. 3 shows the adsorption isotherm of Cs⁺ vs. equilibrium concentration. This isotherm permits distinguishing specific adsorption behaviours. At the lowest concentration, the adsorption of Cs⁺ reaches a plateau, and then decreases at the X₀ site saturation. A change in slope occurs at log₁₀[Cs]ₑ𝑞 ca. -7, which highlights the contribution of the X₁ site at higher total Cs concentration. As in Fig. 2, only two of the exchange sites contributions appear in Fig. 3 because they drive the adsorption at this pH, i.e. pH 7. In term of reactivity, the X₀ site can be compared to the low concentration high affinity frayed edge site (FES), known to be present in illite (Brouwer et al., 1983; Missana et al., 2014b; Poinssot et al., 1999a, b; Sawhney, 1972), and also in micaceous minerals (Zaunbrecher et al., 2015) or vermiculite (Dzene et al., 2015).

At I = 0.01 M (Fig. 2b), the isotherm is overestimated by the modelling. Poinssot et al. (1999a) obtained a similar and unexplained modelling result. Studying the stability of the illite, Poinssot et al. (1999a) have observed a leaching of K⁺ in solution, which could compete with Cs⁺ and lower its adsorption. Bradbury and Baeyens (2005) also supposed that this overestimation could come from the dissolution — leaching of measured Al, Mg, and Ca — of illite and proposed to take into account the behaviour of these pH-leached metals. Applying this proposition requires a measurement of the cationic content in equilibrium solutions and a thermodynamic characterization of the dissolution of clay minerals — see e.g. Gailhanou et al. (2012; 2017) —, which was not available in Poinssot et al. (1999a).
Fig. 2. Adsorption isotherm of Cs vs. pH onto Na-illite at ionic strength 0.1 (a) and 0.01 (b) M (NaClO₄) and [Cs]$_{\text{total}} < 10^{-8}$ M, and modelling with the MSIE formalism. Data digitized from Poinssot et al. (1999a, 1999b). The $X_2$ weak contribution to $\log_{10} R_0$ does not appear in this representation.
Fig. 3. Adsorption isotherm of Cs vs. equilibrium concentration onto an Na-illite at $I = 0.1$ M (NaClO$_4$), pH 7, and modelling with the MSIE formalism. Data digitized from Poinssot et al. (1999a). The $X_2$ and $X_3$ weak contributions to log$_{10} R_D$ at pH 7 do not appear in this representation.

The fitted values for log$_{10} K_{\text{Cs}^+/H^+}^{X_i}$ are summarized in Table 3. These selectivity coefficients — especially associated to the $X_0$ site, which is showing a similar reactivity to FES — can be compared with literature data. Missana et al. (2014a) proposed a log$_{10} K_{\text{Na}^+/\text{Cs}^+}^{\text{FES}} = 6.90$ comparable with the one from Poinssot et al. (1999a, 1999b), i.e. 6.95. In our case, we obtain a similar value with data from Table 3.

$$\log_{10} K_{\text{Na}^+/\text{Cs}^+}^{X_0} - \log_{10} K_{\text{Na}^+/H^+}^{X_0} - \log_{10} K_{\text{Cs}^+/H^+}^{X_0} = 6.59 \pm 1.15$$ (8)

3.2.2.2 Strontium adsorption

The strontium adsorption isotherms onto the purified Na-illite du Puy were extracted from Poinssot et al. (1999b) and Altmann et al. (2015). The selectivity coefficients of the Sr$^{2+}$/2H$^+$ exchange were also determined by simultaneous adjustment of all the adsorption isotherms vs. pH in Fig. 4, and vs. equilibrium concentration in Fig. 5. Missana et al. (2008) also obtained a similar adsorption behaviour vs. pH on an Na-conditioned Silver Hill illite at $I = 0.1$ M (NaClO$_4$). Missana et al. (2008) — original data in NaClO$_4$ provided by the author — obtained a medium log$_{10} R_D$ which is not significantly distinct — ca. 2.1, against ca. 2.0 for Poinssot et al. (1999b) between pH 5 and 8, by considering a log$_{10} R_D$ uncertainty ca. 0.15 log unit. Conversely, it seems that the log$_{10} R_D$ obtained at pH higher than 9 by Missana et al. (2008) were slightly lower than the data from Poinssot et al. (1999b), but close to the ones from Altmann et al. (2015). At $I = 0.01$ M (NaClO$_4$) it also seems that values from Poinssot et al. (1999b) and Altmann et al. (2015) are significantly
lower than the ones obtained by Missana et al. (2008) in neutral pH range — i.e. $\log_{10} R_D$ ca. 3.6 for Poinssot et al. (1999b) and ca. 4.3 for Missana et al. (2008). It was decided nevertheless to principally sit our modelling on the data from Poinssot et al. (1999b) and Altmann et al. (2015) in view of the common illite origin. The obtained modelling will nevertheless be compared with the data from Missana et al. (2008).

Fig. 4 shows the adsorption isotherm of Sr$^{2+}$ vs. pH at $I = 0.1$ (Fig. 4a) and 0.01 (Fig. 4b) M (NaClO$_4$), respectively, and $[\text{Sr}]_{\text{total}} < 10^{-9}$ M. At $I = 0.1$ M, the adsorption is strongly increasing at acidic pH then is reaching a plateau in the pH span 5-8. At more alkaline pH, the adsorption is increasing again. At $I = 0.01$ M, the $\log_{10} R_D$ raised to a plateau at pH below 6 (Fig. 4b). These adsorption behaviours are reminiscent of Sr$^{2+}$ onto Na-smectite (Siroux et al., 2017).

Fig. 5 shows the adsorption isotherm of Sr$^{2+}$ vs. the equilibrium concentration onto an Na-illite at pH 7 and $I = 0.1$ M (NaClO$_4$). Here, the adsorption stays constant, i.e. $\log_{10} R_D = 2$, then decreases at the highest concentrations starting from $[\text{Sr}]_{\text{eq}} = 10^{-5}$ M. Once again, this adsorption behaviour recalls the Sr one onto Na-smectite (Siroux et al., 2017). Therefore, it is possible to state that Sr$^{2+}$ has similar adsorption behaviours onto Na-illite and -smectite when it competes with Na$^+$. The modelling represents well the experimental data except for the pH isotherm at $I = 0.01$ M. Actually, the isotherms in Fig. 4a and Fig. 5 were acquired at the same ionic strength (0.1 M), and weigh heavily on the corrected coefficient values compared to the data from the isotherm at $I = 0.01$ M. Consequently, the modelling outcome overestimates the data of the experiment at $I = 0.01$ M. The reasons of this overestimation are probably the same that in § 3.2.2.1. The obtained $\log_{10} K_{\text{Sr}^{2+}/2\text{H}^+}^{X_i}$ are reported in Table 3. As supposed in this study, $\log_{10} K_{\text{Sr}^{2+}/2\text{H}^+}^{X_1}$ for illite and $\log_{10} K_{\text{Sr}^{2+}/2\text{H}^+}^{Y_1}$ for smectite are similar except for the $X_1$ and $Y_1$ sites. It seems that the $X_0$ site does not have any significant impact in the Sr$^{2+}$ adsorption modelling, and it was decided not to consider it in a first approximation.

Comparison with data from Missana et al. (2008) is also shown in Fig. 4. As awaited from the similarity of $\log_{10} R_D$ values in the neutral pH range, the data from Missana et al. (2008) are well represented by the modelling. One can notice that the 0.01 M NaClO$_4$ pH-isotherm is better described than the data from Poinssot et al. (1999b) and Altmann et al. (2015), which are used to fit the selectivity coefficients. Nevertheless, the agreement between experiments and simulation is clearly less satisfying at extremum pH values. The reasons behind these differences are not clear and may be a consequence of the clay mineral dissolution as reminded earlier.
Fig. 4. Adsorption isotherm vs. pH of Sr$^{2+}$ onto an Na-illite at I = 0.1 M (NaClO$_4$) (a), and I = 0.01 M (NaClO$_4$) (b), [Sr]$^{\text{total}} < 10^{-9}$ M and S/L ratio = 1.78 g.L$^{-1}$ and modelling with the MSIE formalism. Data digitized from Poinssot et al. (1999b) (diamonds) and Altmann et al. (2015) (triangles), and from Missana et al. (2008) (circles).
Fig. 5. Adsorption isotherms in concentration of Sr\(^{2+}\) onto a Na-illite at I = 0.1 M (NaClO\(_4\)), pH 7 (a) and pH 8.3 (b) and S/L ratio = 1.75 g.L\(^{-1}\) and modelling with the MSIE formalism. Data digitized from (a) Poinssot et al. (1999b) and (b) Altmann et al. (2015). The X\(_3\) weak contribution to log\(_{10}\)RD at these pH values do not appear in this representation.

### 3.3 Verification of the Additivity Approach of the MSIE Modelling

To apply the MSIE modelling with the database in more complex media as soils or sediments, it is possible to apply the additivity approach by combining the relative contribution of each reactive phases in the global adsorption of the mixture. In this study, the additivity approach was used in order to check the efficiency and robustness of this
database to simulate the adsorption of Cs\(^+\) and Sr\(^{2+}\) onto illite/smectite mixtures with MSIE formalism. To achieve this aim, adsorption isotherms onto mixtures of pure phases were acquired. Blind simulations were compared to experimental data, by using the \(\log_{10} K_{Mn^+/nH^+}^{*X_i}\) for illite, and \(\log_{10} K_{Mn^+/nH^+}^{*X_j}\) for smectite from Siroux et al. (2017) (cf. Table 3).

### 3.3.1 Adsorption of Cs\(^+\) onto Na-illite/Na-smectite mixtures

The adsorption isotherms of Cs\(^+\) on the two mixtures at \(\text{pH} 7, I = 0.1 \, \text{mol.kg}^{-1} \, \text{NaCl}, \) and \(S/L = 3.8 \, \text{g.L}^{-1}\) is shown in Fig. 6. As in § 3.2.2.1, the isotherm follows the specific adsorption of Cs\(^+\) onto illite, which seems to mainly drive the adsorption. Despite the higher percentage of smectite in M1 than in M2, both adsorption behaviours of Cs\(^+\) onto the mixtures are almost identical. Smectite seems to drive the adsorption at higher total concentration of Cs for M1 — \(\log_{10}([\text{Cs}]_{\text{eq}} /\text{mol.kg}^{-1}) > -4\) —, and for M2 smectite equals illite contribution only at the highest total Cs concentration — \(\log_{10}([\text{Cs}]_{\text{eq}} /\text{mol.kg}^{-1})\) ca. –3. The shape of the Cs\(^+\) adsorption isotherm is well reproduced by the simulation, although experimental results are slightly overestimated at low Cs concentration. This overestimation is especially higher for M2, which contains the higher illite proportion. The origin of these overestimations will be further questioned in the following.

The uncertainty of the simulation can be calculated from the uncertainties on the available parameters in Table 3 using propagation of errors — uncertainties on \(X_0\) site capacity and \(\log_{10} K_{\text{Na}^+/H^+}^{*X_0}\) for illite are not accounted for. It can be seen (Fig. 6) that at least from the point of view of the fitted parameters uncertainties, the overestimation of the adsorption of Cs\(^+\) onto the clay minerals mixtures cannot be considered significant. As awaited from values in Table 3, the major contributor to the uncertainty being the selectivity coefficient of the Na\(^+\)/H\(^+\) exchange on the high affinity site of illite \(X_0\), \(\log_{10} K_{\text{Na}^+/H^+}^{*X_0}\), at trace concentration of Cs\(^+\) — affecting a 100\% uncertainty on \(X_0\) site capacity and \(\log_{10} K_{\text{Cs}^+/H^+}^{*X_2}\) does not greatly affect the global feature.
3.3.2 Adsorption of Sr\textsuperscript{2+} onto Na-illite/Na-smectite mixtures

The adsorption isotherms of Sr\textsuperscript{2+} onto both mixtures are presented in Fig. 7. As discussed in § 3.2.2.2, smectite and illite are showing comparable Sr\textsuperscript{2+} adsorption behaviours, which
can explain that the obtained log_{10}R_D are similar for both mixtures. For Sr^{2+}, the adsorption is both linked to the proportion of each exchanger — i.e., in mixture M1 or M2 — and the selectivity coefficients for the 2Na^{+}/Sr^{2+} exchange, log_{10}K_{2Na^{+}/Sr^{2+}}^{x_i}.

If one considers only the exchange capacity, the adsorption of Sr^{2+} should be slightly higher in mixtures with higher concentration of smectite. This is shown in Fig. 7, where the adsorption of Sr^{2+} is slightly higher onto M1 (Fig. 7a) than onto M2 (Fig. 7b) at log_{10}([Sr]_eq /mol.kg\textsuperscript{-1}) > -3.3. However, this seems difficult to assess as the final log_{10}R_D values for the two mixtures are almost the same — i.e. log_{10}R_D are 2.11 and 2.01 for M1 and M2 mixtures, respectively. Despite a very slight overestimation, in particular for the M2 mixture, the agreement between simulations and experimental results is good. Simulations also follow well the decrease in adsorption caused by the addition of illite. It seems that in the case of Sr^{2+}, the database used allows a correct representation of the adsorption behaviour onto illite/smectite mixtures.

This non-significant difference between M1 and M2 isotherms is coming from the higher selectivity coefficients for the 2Na^{+}/Sr^{2+} exchange on Na-illite,

\[ \log_{10}K_{2Na^{+}/Sr^{2+}}^{x_i} = 2 \log_{10}K_{Na^{+}/H^{+}}^{x_i} - \log_{10}K_{Sr^{2+}/2H^{+}}^{x_i} \]

the values of which are shown in Table 4 — calculated from Table 3. The selectivity coefficients for illite are at least an order of magnitude higher than the one for smectite, when the total site capacity is 3.6 times higher for smectite compared to illite. Considering, pure Na-illite (this study) and pure Na-smectite (Siroux et al., 2017), similar R_D at 0.1M NaClO_4 are observed. Consequently, combining both minerals in various proportions leads to comparable R_D values.

**Table 4.** Total site capacities of the clay minerals, and selectivity coefficients for the 2Na^{+}/Sr^{2+} exchange calculated using values from Table 3.

<table>
<thead>
<tr>
<th>Clay Type</th>
<th>Capacity (mmol.100g\textsuperscript{-1})</th>
<th>( \log_{10}K_{2Na^{+}/Sr^{2+}}^{x_1} )</th>
<th>( \log_{10}K_{2Na^{+}/Sr^{2+}}^{x_2} )</th>
<th>( \log_{10}K_{2Na^{+}/Sr^{2+}}^{x_3} )</th>
</tr>
</thead>
<tbody>
<tr>
<td>Illite</td>
<td>(24.9 ± 1.9)</td>
<td>1.79 ± 0.16</td>
<td>2.38 ± 0.50</td>
<td>3.46 ± 0.38</td>
</tr>
<tr>
<td>Smectite</td>
<td>(88.7 ± 12.1)</td>
<td>0.68 ± 0.08</td>
<td>1.00 ± 0.19</td>
<td>2.91 ± 0.28</td>
</tr>
</tbody>
</table>

As in the case of Cs\textsuperscript{+} adsorption on the clay minerals mixtures, the uncertainties of the Sr\textsuperscript{2+} simulations can be calculated by propagation of errors using the values in Table 3. It can be seen (Fig. 7) that the simulations are not significantly different. Nevertheless, the origin
of the uncertainties are slightly different. The main contributors to the uncertainty being: $\log_{10}K_{\text{Na}^+/H^+}^{Y_1}$, $\log_{10}K_{\text{Na}^+/H^+}^{Y_2}$, and $\log_{10}K_{\text{Sr}^{2+/2H^+}}^{Y_2}$ from smectite in the case of M1 mixture; and $\log_{10}K_{\text{Na}^+/H^+}^{X_1}$ and $\log_{10}K_{\text{Na}^+/H^+}^{X_2}$ from illite in the case of M2 mixture.

Fig. 7. Adsorption isotherms of Sr vs. equilibrium concentration at $I = 0.1$ mol.kg$^{-1}$ (NaCl), pH 7 and S/L ratio = 3.8 g.L$^{-1}$ on M1 (a) and M2 (b) — see composition in Table 2 — with the associated simulations (plain) and uncertainty (1$\sigma$, dot), compared with illite (dash) and smectite (dash-dot) components.
3.3.3 Discussion on the additivity results

The measured CEC of the two purified clay minerals used in this study are showing different values compared to those from the database: smectite has a CEC of 88.7 meq.100g⁻¹, against 83.3 meq.100g⁻¹ for our SWy-2 smectite, and illite has a CEC of 24.9 meq.100g⁻¹, against 20.9 meq.100g⁻¹ for our illite du Puy. Given the XRD analysis (cf. § 2.1), one can postulate that the used smectite and illite contained impurities. In view of the measured CEC, it is possible to estimate that the smectite and illite are 94% and 84% pure, respectively. These impurities produce a lesser impact on the sample CEC than the uncertainty calculated on the total site capacity — cf. Table 3 for smectite (88.7 ± 12.1) mmol.100g⁻¹ or ca. 14%, and illite (24.8 ± 1.9) mmol.100g⁻¹ or ca. 8% only accounting for X₁ to X₃ site. As seen earlier, these differences may be integrated through the uncertainties of the simulations. This impurity rates are accounted for in Fig. 8, implying a lowering of log₁₀R₀. Although the adsorption of Cs⁺ onto M1 (Fig. 8a) is better represented, the representation of M2 experiment is still slightly overestimated. In the case of Sr²⁺ (Fig. 8b), the impurity of clay minerals does not significantly change the simulations. In all cases the account of purity is still within the uncertainties of the simulations (vide ante).

The impurity of the clay minerals is not sufficient to explain the slight overestimation of adsorbed Cs onto the clay mixtures. Nevertheless, this problem could come from the clay minerals themselves, especially the illite du Puy. As seen earlier (cf. § 2.1), the purified and Na-conditioned illite du Puy still presents traces of exchangeable K⁺, i.e. 1.1 meq.100 g⁻¹, which represents around 5% of the illite CEC. As Cs⁺, K⁺ is well known to interact strongly with illite — especially with the FES —, the presence of K⁺ could interfere with the adsorption of Cs⁺ even at trace concentration. Missana et al. (2014a) showed this effect studying the adsorption of Cs⁺ onto various conditioned illite. Onto K-illite the adsorption of Cs⁺ was shown to be lower than onto Na-illite — a decrease of around 2.2 units of log R₀ was observed at low Cs concentration.
Fig. 8. Adsorption isotherms vs. equilibrium concentration of Cs (a) and Sr (b) at $I = 0.1$ mol.kg$^{-1}$ (NaCl), pH 7 and S/L ratio = 3.8 g.L$^{-1}$ on mixture M1 (squares) and M2 (diamonds) (composition in Table 2); the associated simulations for M1 in Fig. 6 (plain) and M2 in Fig. 7 (dash) are compared to the simulation taking into account of the clay minerals purity for M1 (dot) and M2 (dash-dot).
3.3.4 Discussions on the limitations and advantages of the MSIE modelling

Several limitations are to be considered for this kind of modelling exercise. As already emphasized, applying selectivity coefficients that are fitted even at different ionic strengths to whatever ionic strength could be a problem if the strong hypothesis of the constant ratio of the activity coefficient of adsorbed species is not verified. The same applies if the change in ionic strength is inducing structural modification within the clay minerals. Considering that an illite or a smectite from sedimentary locations, from a soil, or even synthetic clays, are showing the same adsorption properties is also a strong hypothesis, which can induce caution in the outcome of the modelling.

The parameters can be applied in various conditions. First, it applies in wide ranges of physico-chemical parameters, i.e. ionic strength and pH, which ensures its applicability in various contexts. Second, it is applicable at various mass upon volume ratios, which ensures its applicability in reactive transport in porous media. Finally, any competition adsorption reaction, or side reaction in solution, can be taken into account as long as the selectivity coefficient, or thermodynamic constant, are available.

4 CONCLUSIONS

The adsorption of Cs\(^+\) and Sr\(^{2+}\) onto a purified Na-illite has been studied and compiled from the literature. These isotherms were modelled within the MSIE formalism with the aim of obtaining corrected selectivity coefficients for the Na\(^+\)/H\(^+\), Cs\(^+\)/H\(^+\), and Sr\(^{2+}\)/2H\(^+\) ion-exchange reactions — Na\(^+\)/Cs\(^+\) and Sr\(^{2+}\)/2Na\(^+\) can be calculated from these values. The experiments are generally well represented by the modelling. These representations assure a good efficiency and robustness of the obtained database to simulate the adsorption of Cs\(^+\) and Sr\(^{2+}\) onto Na-illite.

Afterwards, the performance of the modelling and the database, including the parameters from Siroux et al. (2017), was verified onto illite/smectite mixtures. Despite overestimations observed at low Cs concentration, the good agreement between the experimental isotherms of the adsorption of Cs\(^+\) and Sr\(^{2+}\) with the simulations tends to confirm the robustness of the database, which could be used to predict the adsorption behaviour of Cs\(^+\) and Sr\(^{2+}\) onto more complex phases as sediments or soils.

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